## Castleconnell GWB: Summary of Initial Characterisation.

Hydrometric Area			Associated surface water features	Associated terrestrial	Area (km <sup>2</sup> )	
25 - Mulkear/ Lower Shannon catchment		hannon catchment	Rivers: Shannon, Mulkear.		25	
Topography	This narrow GWE and flat-lying to g of the Mulkear to volcanic rocks. Dr rocks.	3 is elongated in N-S of ently undulating. Gro 81 mAOD at Knockb rainage density is low	lirection, and is about 9.5 km long by 3 km wide. In general, the ground is 20–40 mAOD, und elevation within the GWB ranges from 20 mAOD along the Shannon and lower reaches rack. The highest ground occurs in the east of the GWB, on higher ground underlain by over most of the GWB, but higher just south of the centre of the GWB, near to the Volcanic			
	Aquifer categories	The GWB comprises predominantly an <b>Rk<sup>d</sup></b> : Regionally important karstified aquifer dominated by diffuse flow. An area about 2km <sup>2</sup> of volcanic rock occurs in the east of the GWB; it is currently classified as an <b>Lm:</b> Locally important aquifer which is generally moderately productive.				
eology and Aquifers.	lithologies	Rocks in the centre-east of the GWB.				
	Key structures	The rocks of this GWB occur on the SE limb of a WSW-plunging syncline. Bedding dips generally NE at low angles of approximately 10-20°. There are minor folds mapped in the northern part of the GWB. Fractures and joints may be more open on the axis of the minor anticlines. Conversely, fractures may be closed in the minor synclines and in the north of the GWB, which is near the major fold axis.				
	Key properties	Transmissivities in diffusely karstified aquifers are in the range 20–2000 m <sup>2</sup> /d. In this area of the country, the median value will probably be towards the lower-middle end of the range. At Croom and Fedamore WSs in the nearby Fedamore GWB, transmissivities are $120 \text{ m}^2/\text{d}$ [estimate range $95-145 \text{ m}^2/\text{d}$ ] and $34 \text{ m}^2/\text{d}$ [estimate range $23-41 \text{ m}^2/\text{d}$ ], respectively. Groundwater gradients within the karstic aquifer are low, ranging from approximately 0.005 to 0.01. Storativity in this aquifer is low (effective porosity ~1.5-2.5%). Transmissivities in the volcanic rocks are variable; in places, clays from weathering after their deposition have blocked the fissures; in other areas, these weathering products have been washed out of the fracture system. Transmissivities will be in the range $2-100 \text{ m}^2/\text{d}$ . Groundwater gradients can be up to 0.05 in these rocks, since they tend to form higher ground with steep slopes.				
9		(data sources: Rock estimates from map:	Unit Group Aquifer Chapters, Limerick ( s)	<i>GWPS and Source Reports, see reference</i>	<i>:es;</i>	
	Thickness	The Dinantian Pure Unbedded Limestones attain maximum thicknesses of more than 1200 m. However, most groundwater flows in an epikarstic layer a few metres thick and in a zone of interconnected solutionally-enlarged fissures and conduits extending about 30 m below this. Deeper groundwater flows can occur along fault zones. On Aughinish island, on the south side of the Shannon Estuary, there are very deep (~ 60 mbsl) conduits that relate to an ancient baselevel. There may be such conduits in this area, but they are not known; field work would be required to confirm their presence or otherwise. Two boreholes in the adjacent GWB with low yields and big drawdowns had inflows between around 75 mbgl and 115 mbgl. The low yields indicate flow through small fissures, not conduits. Within the volcanic rocks, most groundwater flux is likely to be in the top $\leq 20$ m, this zone comprising a weathered layer of a few metres and a connected fractured zone below this, although more isolated water-bearing fractures or faults can be intercepted at greater depths.				
lying Strata	Lithologies	The GWB is mainly Alluvium within the eastern side, there a of the Gardenhill an	<ul> <li>covered by Limestone Till subsoils. Ther</li> <li>Till that are separately mapped. At the fc</li> <li>re Peat deposits, together with a complex</li> <li>id Gooig Gravel GWBs overlie the NE of</li> </ul>	ce are small pods of gravel and of Undif oot of the hill underlain by volcanic rocl assemblage of Alluvium, and Limeston this GWB.	fferentiated k, on the le Till. Parts	
	Thickness	There are few data a and south of the GB indicate subsoils thi	vailable for this area. Rock outcrops are s W, excepting along the River Shannon in cknesses in the range 7-24 m.	small and scattered, occurring only in the north of the GWB. Available thicks	ie centre ness data	
Ovei	% area aquifer near surface	[Information to be a	ıdded at a later date]			
	Vulnerability	Groundwater vulner areas where rock ou	ability over most of the GWB is High. The trops. In the NE of the GWB, there is an	ere are small areas of Extreme vulneral area of Moderate vulnerability.	bility in the	
Recharge	Main recharge mechanisms	Diffuse recharge wi aquifer via outcrop. over most of this G However, in low-lyi rejected recharge, i. Linear recharge may low transmissivity C	Il occur over most of the GWB via rainfal The epikarst redistributes diffuse recharge WB indicates that potential recharge readi ing areas with a high water table in this hi e. a proportion of the effective rainfall is r y occur along losing river stretches where GWB to the south and east.	I soaking through the subsoil and direct e in the subsurface. The lack of surface ly percolates into the groundwater syste ghly transmissive aquifer, there may be rejected due to lack of storage space in the the Mulkear River crosses onto this GV	tly to the drainage em. some the aquifer. WB from the	
	Est. recharge rates	[Information to be a	idded at a later date]		1.000000000000000000000000000000000000	

arge	Important springs and high yielding wells (m <sup>3</sup> /d) Main discharge mechanisms	There are no High yielding springs (>2,160 m <sup>3</sup> /d) or Excellent (>400 m <sup>3</sup> /d) wells known in this GWB. Near Annacotty, Limerick Horsepackers Ltd. abstract 136 m <sup>3</sup> /d from a 57 m deep borehole, indicating at least a Good (100 m <sup>3</sup> /d < yield < 400 m <sup>3</sup> /d) yield. The main discharges are to the streams crossing the aquifer, and to the River Shannon in the NW of the GWB and the Mulkear River in the south of the GWB. There is probably some cross-flow from this GWB to the lower
schi		transmissivity pure bedded limestone aquifer of the Limerick East GWB.
Ξ.	Hydrochemical Signature	There are no data available to assess this GWB. The hydrochemistry of groundwaters from the nearby Fedamore GWB indicates Very Hard (370–430 mg/l as CaCO <sub>3</sub> ), calcium-bicarbonate type waters with high alkalinities (330–380 mg/l as CaCO <sub>3</sub> ) and electrical conductivities, and neutral pHs. Conductivities range between 720–900+ $\mu$ S/cm. In general, background chloride concentrations will be higher than in the Midlands, due to proximity to the sea.
Gro	undwater Flow Paths	The limestone rocks are devoid of intergranular permeability; groundwater flows through a diffuse network of solutionally-enlarged fissures and small conduits, and along faults. Dissolutional enlargement of joint, fracture and fault planes is the major mechanism that has created permeability. Groundwater is likely to flow in two main hydrogeological regimes: (1) an upper, shallow, highly karstified weathered zone, known as the epikarst, in which groundwater moves quickly, through solutionally enlarged conduits, in rapid response to recharge; (2) a deeper zone, where a dispersed slow groundwater flow component in smaller fractures and joints outside the main conduit systems. In some areas, the aquifer may be highly karstified, with groundwater flowing through interconnected, solutionally enlarged conduits and cave systems. These localised high permeability zones can give rise to rapid groundwater velocities. The GWB is considered to be unconfined. It is considered that the rivers and streams are in hydraulic continuity with the aquifer. Therefore, they represent the water table elevation. Near streams and rivers, water levels should be within 2 m of ground level. Figure 1 shows a hydrograph from a borehole near Annacotty that is situated near to the Mulkear River. The seasonal water level variation is less than 1 m. In relatively elevated areas between rivers, the water table may be considerably deeper, giving significant unsaturated zones. The water table is likely to generally follow the topography. Groundwater flow paths in this GWB are generally long, and can be up to several km's long. Groundwater may also discharge locally to streams, however, if the topography is variable. In discharge zones, flow paths will be shorter, around 100–300 m. The regional groundwater flow direction in the norther half of the GWB is from south to north, to the River Shannon. In the south of the GWB, groundwater flows at an oblique angle to the Mulkear River. Groundwater may also flow westwards into the Limerick City East GWB. L
		or alteration minerals occur, this may block the flow conduits, unless they have been flushed from the system by high groundwater gradients in the hilly terrain. In the zones where the fractures and joints are not clogged with clays, transmissivities can be relatively high. Groundwater is unconfined; the water table follows the topography and significant unsaturated zones occur in elevated areas. In general, flows in the aquifer are likely to be concentrated in a thin zone near the top of the rock; the weathered zone may be up to 3 m thick, with a connected fractured zone a further 15-20 m, below which is a generally poorly fractured zone
Gi	roundwater &	There is an effective hydraulic interconnection between groundwater and surface water in the karst limestone.
S	urface water	Groundwater is discharged to the surface as baseflow to streams and rivers crossing the GWB, and to the River
i	interactions	Shannon at the NW of the GWB.

Conceptual model	<ul> <li>The terrain GWB at the boundary are formed with the P</li> <li>The GWE conduits a been enlar zones hav and faults</li> <li>Recharge the Mulke rejected in</li> <li>Groundwa comprises conduits, flow may in the adja</li> <li>It is possil thus affec</li> <li>Most of the significan whereas, in</li> <li>Groundwa GWB, with in the northengths ar</li> <li>Rivers matical structure significan whereas and the significan whereas and the significan whereas are significant.</li> </ul>	WB at the hill underlain by volcanic rocks. It is bounded to the NW by the River Shannon. The northern, eastern and southern ooundary is formed by the contact with the lower transmissivity rocks of the Lough Graney GWB. The southern/SE boundaries re formed by the contact with the low transmissivity Lower Impure limestones of the Slieve Phelim GWB. The contact with the vith the Pure Bedded Limestones of the Limerick City East GWB form the western marigin. The GWB mainly comprises diffusely karstified limestones in which groundwater is transmitted through a network of small onduits and fissures, and an epikarstic zone. The fault and fracture network and bedding-parallel pressure solution planes have eee enlarged by dissolution, resulting in a highly transmissive aquifer with rapid groundwater flow in which the more permeable tones have specific orientations. In the relatively lower transmissivity volcanic rocks, groundwater flow occurs in fractures, joints and faults. The aquifers have low storativity. Recharge occurs diffusely through the subsoils or at rock outcrop. Linear recharge may occur along losing river stretches where he Mulkear River crosses onto this GWB from the low transmissivity GWB to the south and east. Potential recharge may be ejected in areas where the water table is very close to the surface. Groundwater flux in this aquifer will be concentrated in an approximately 30 m zone at the top of the bedrock. This zone comprises an epikarstic layer of a few metres, below which is a network of diffuse solutionally-enlarged joints and small onduits, fractures and faults. Deeper groundwater flow can cocur along permeable fault zones or deeper fractures. The volume of low may be significantly less in these deeper systems, as evidenced by the poor yields and drawdown characteristics of boreholes n the adjacent GWB in which inflows were measured at around 75-115 mbgl. t is possible that, as the focus of drainage, the area near to the River Shannon is more karstified than other parts of thi			
Attac	hmonts	Groundwater hydrograph (Figure 1)			
Instru	imentation	EPA Water Level Monitoring boreholes: Annacotty (LIM 165)			
Inform	nation	Deakin L and Daly D (2000) County Clare Groundwater Protection Scheme Geological Survey of Ireland Report to			
Sourc	es	Clare Co. Co., 67 pp.			
		Deakin, J., Daly, D. and Coxon, C. (1998) County Limerick Groundwater Protection Scheme. Geological Survey of			
		Ireland Report to Limerick Co. Co., 72 pp.			
		Deakin, J. (1995) Croom WS – Groundwater Source Protection Zones. Geological Survey of Ireland Report to			
		Limerick Co. Co., 6 pp.			
		Deakin, J. (1995) Fedamore WS – Groundwater Source Protection Zones. Geological Survey of Ireland Report to			
		Limenck Co. Co., 6 pp.			
D: .	•	Aquiter chapters. Dinanuan Pure Unbedded Limestones.			
Discla	imer	Note that all calculations and interpretations presented in this report represent estimations based on the information			
		sources described above and established hydrogeological formulae			



Figure 1: Groundwater hydrograph



Rock units in GWB

Rock unit name and code	Description	Rock unit group
Waulsortian Limestones (WA)	Massive unbedded lime-mudstone	Dinantian Pure Unbedded Limestones
Tuff (Tu)		Basalts and other Volcanic rocks